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# Isochron Dating Method

FOR

# DUMMIES<sup>®</sup>

*A Reference  
for the  
Rest of Us!*

Radiometric  
dating  
explained in few  
words

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This text follows the structure of article by Young (2007). Radioactive decay is an exponential process. That is, the rate of decay of isotope  $N$  is proportional to its quantity:

$$\frac{dN}{dt} = -\lambda N, \quad (1)$$

where  $\lambda$  is the decay constant, related to the half-life of radioactive isotope as  $\tau_{1/2} = \ln 2/\lambda$ . Isotope half-lives are usually determined experimentally and  $\lambda$  values for different atoms and isotopes are available from various sources. Aleksandrov et al. (1963) teach how differential equation 1 can be solved in order to derive an equation that allows to determine the exact quantity of radioactive isotope at any point of time  $t$ :

$$N = N_0 e^{-\lambda t}, \quad (2)$$

where  $N_0$  is the initial quantity of isotope contained in the rock at the time of its formation (crystallization).

A rock sample is considered a closed system if neither parent nor daughter isotopes have been introduced after its formation by an external process. This is a requirement for accurate dating. A closed system can, thus, be characterized as having its:

$$N + D = N_0 + D_0, \quad (3)$$

where  $D$  and  $D_0$  are correspondingly the present-time and initial amounts of the daughter isotope. Combining equations 2 and 3, we obtain equation for radiometric dating:

$$D = D_0 + N(e^{\lambda t} - 1). \quad (4)$$

It is difficult to measure absolute amounts of isotopes in a sample, but mass spectrometers can measure isotope ratios. Thus, we can introduce stable non-radiogenic isotope  $d$  into equation 4:

$$D/d = D_0/d_0 + (N/d)(e^{\lambda t} - 1). \quad (5)$$

With rubidium-strontium dating, whereby  $^{87}\text{Rb}$  decays into  $^{87}\text{Sr}$ , this becomes:

$$^{87}\text{Sr}/^{86}\text{Sr} = (^{87}\text{Sr}/^{86}\text{Sr})_0 + (^{87}\text{Rb}/^{86}\text{Sr})(e^{\lambda t} - 1), \quad (6)$$

where stable isotope is  $^{86}\text{Sr}$ . This equation can be rearranged into

$$^{87}\text{Sr}/^{86}\text{Sr} = (e^{\lambda t} - 1)(^{87}\text{Rb}/^{86}\text{Sr}) + (^{87}\text{Sr}/^{86}\text{Sr})_0, \quad (7)$$

which has a form of linear equation with two variables,  $y = mx + b$ , with slope  $m$  represented by  $(e^{\lambda t} - 1)$ . The slope of the line, thus, is related to time  $t$  passed since crystallization. If  $^{87}\text{Sr}/^{86}\text{Sr}$  values are used as ordinate ( $y$ -axis) and  $^{87}\text{Rb}/^{86}\text{Sr}$  for abscissa ( $x$ -axis), then the constant  $(^{87}\text{Sr}/^{86}\text{Sr})_0$  would correspond to  $b$ , which is the  $y$ -intercept, given that initial and present-time ratios of stable isotope  $^{86}\text{Sr}$  remain constant.

By analyzing multiple samples of isotope ratios ( $^{87}\text{Sr}/^{86}\text{Sr}$  and  $^{87}\text{Rb}/^{86}\text{Sr}$ ) within the same rock, it is possible to plot multiple points on the isochron diagram and determine the slope of the line,  $(e^{\lambda t} - 1)$ —see figure 1. From the slope, knowing the half-time of the isotope (or decay constant  $\lambda$ ), it is possible to find  $t$  of rock crystallization. Initial ratios of isotopes in the rock can also be derived from the  $y$ -intercept.

Bouvier et al. (2015) report that the age of the Moama meteorite is  $4519 \pm 34$  Myr by constructing an isochron diagram for the  $^{147}\text{Sm}-^{143}\text{Nd}$ . Below we perform our own analysis using the equation presented above. The radioactive decay that is used in this case is the alpha decay of  $^{147}\text{Sm}$  to form the daughter  $^{143}\text{Nd}$  with a half life 106 Byr. Table 1 shows the data from Bouvier et al. (2015) for the

## The Rubidium-Strontium System

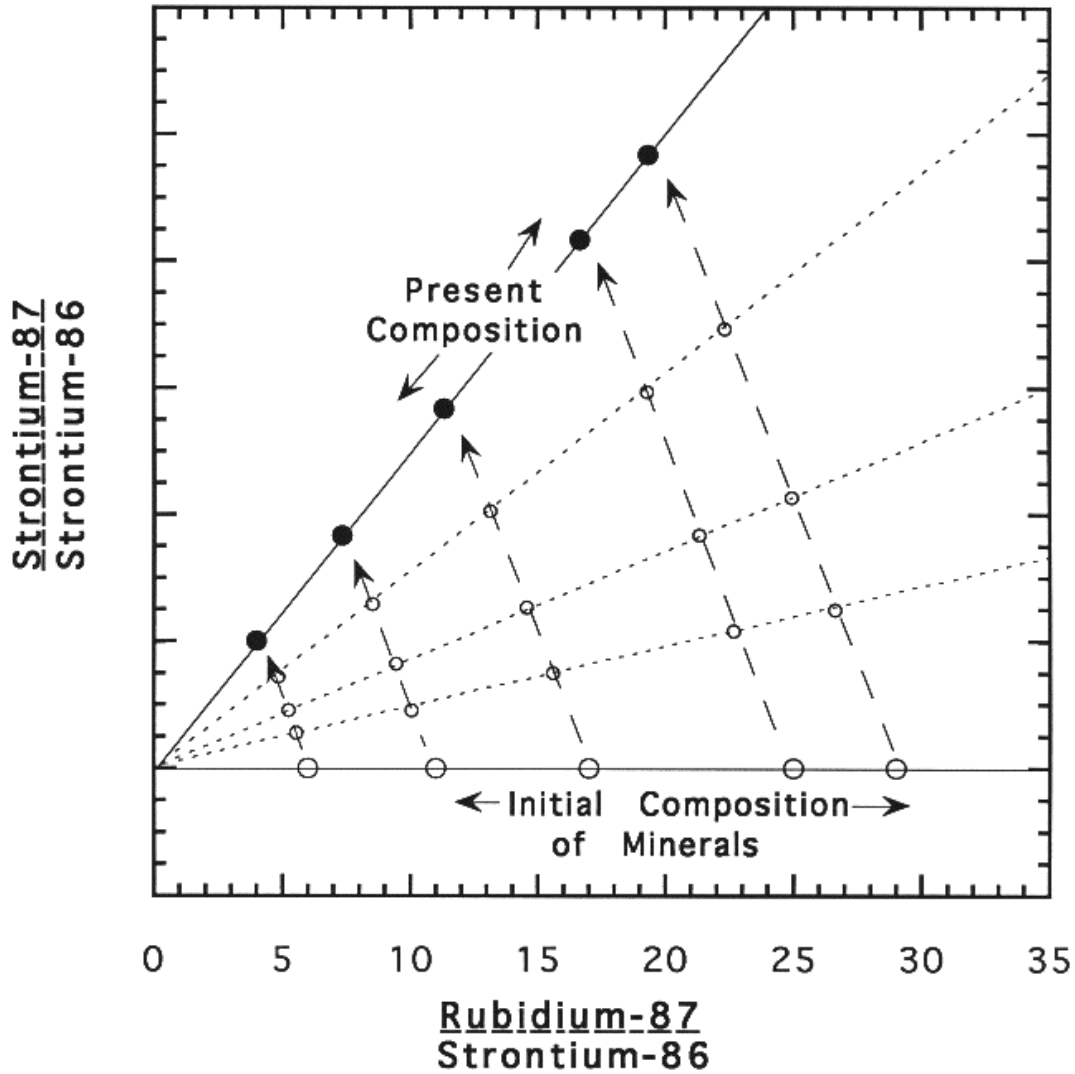


FIG. 1.— Rubidium-strontium isochron diagram.  $^{87}\text{Rb}$  decays into  $^{87}\text{Sr}$ , while stable isotope  $^{86}\text{Sr}$  remains constant. The slope of the isochron is  $(e^{\lambda t} - 1)$ . As the rock becomes older, the isochron pivots around the  $y$ -intercept that corresponds to the initial amounts of isotopes  $(^{87}\text{Sr}/^{86}\text{Sr})_0$ . The older the rock, the steeper is the isochron. Figure by R. Wiens.

pyroxene, plagioclase and whole rock abundances of the  $^{147}\text{Sm} / ^{144}\text{Nd}$  and  $^{143}\text{Nd} / ^{144}\text{Nd}$  ratios. Plotting these ratios gives the chart shown on figure 2. Using the least squares linear regression functions in Excel we determined that this line is  $y = mx + b$ , where  $m = 0.029943 \pm 0.00014$  and  $b = 0.506747 \pm 0.000036$ . From equations above, the slope of the line is given by the expression

$$m = e^{\lambda t} - 1, \quad (8)$$

so

$$t = \frac{\ln(m + 1)}{\ln 2} \times \tau_{1/2} = \frac{\ln(0.029943 + 1)}{\ln 2} \times 106 \text{ Byr} = 4.512 \text{ Byr}. \quad (9)$$

TABLE 1  
MOAMA METEORITE, DATA FROM BOUVIER ET AL. (2015).

Group	$^{147}\text{Sm} / ^{144}\text{Nd}$	$^{143}\text{Nd} / ^{144}\text{Nd}$
Whole rock	0.2317	0.513705
Pyroxene	0.3523	0.517287
Plagioclase	0.1151	0.510184

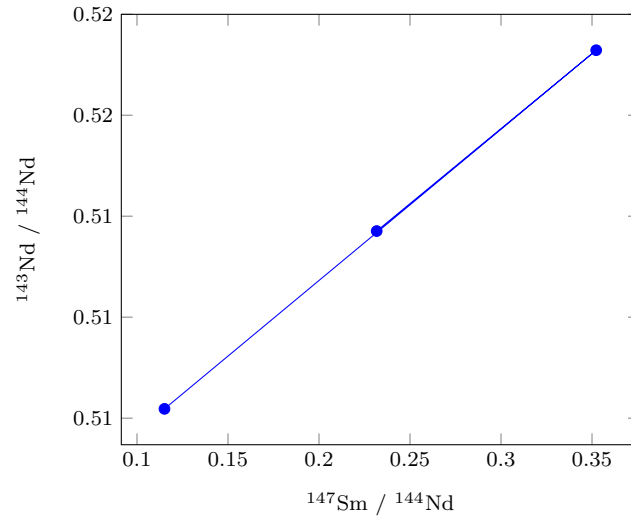


FIG. 2.— Isochron diagram for pyroxene, plagioclase and whole rock ratios of  $^{147}\text{Sm}$ – $^{143}\text{Nd}$  radioactive decay in the Moama meteorite.

This is quite close to the value reported in the paper of 4519 Myr.

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